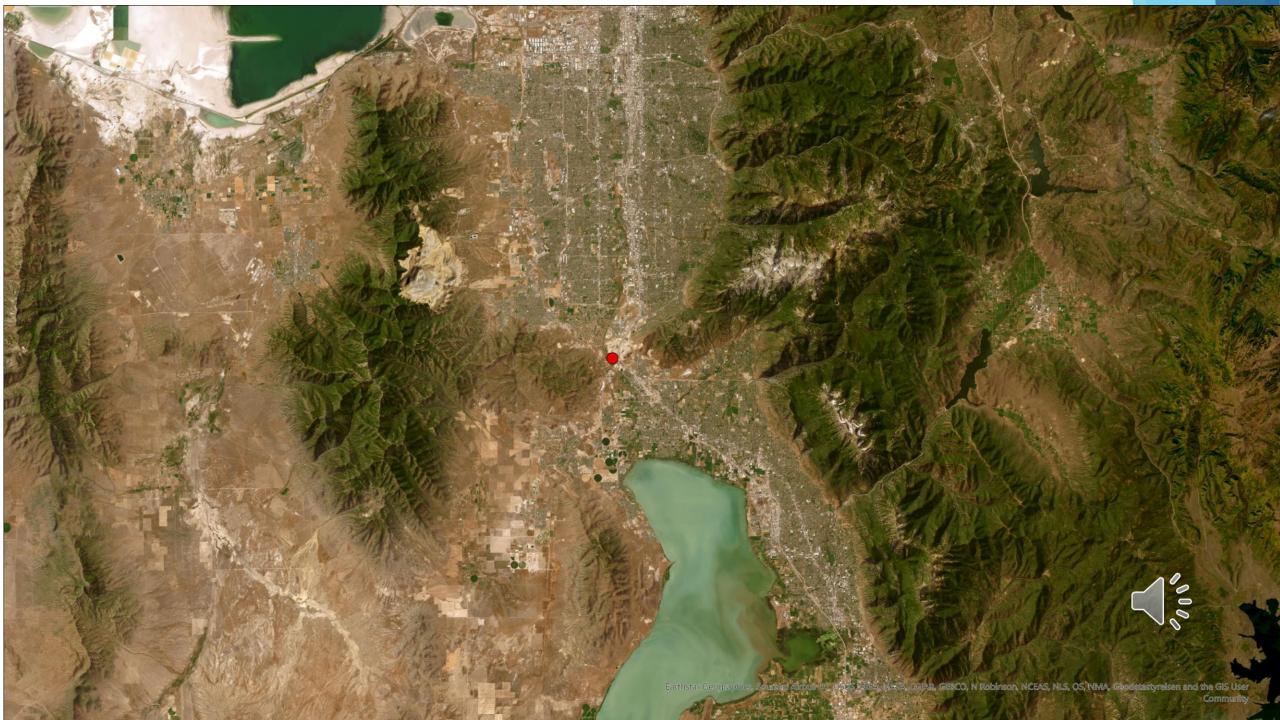
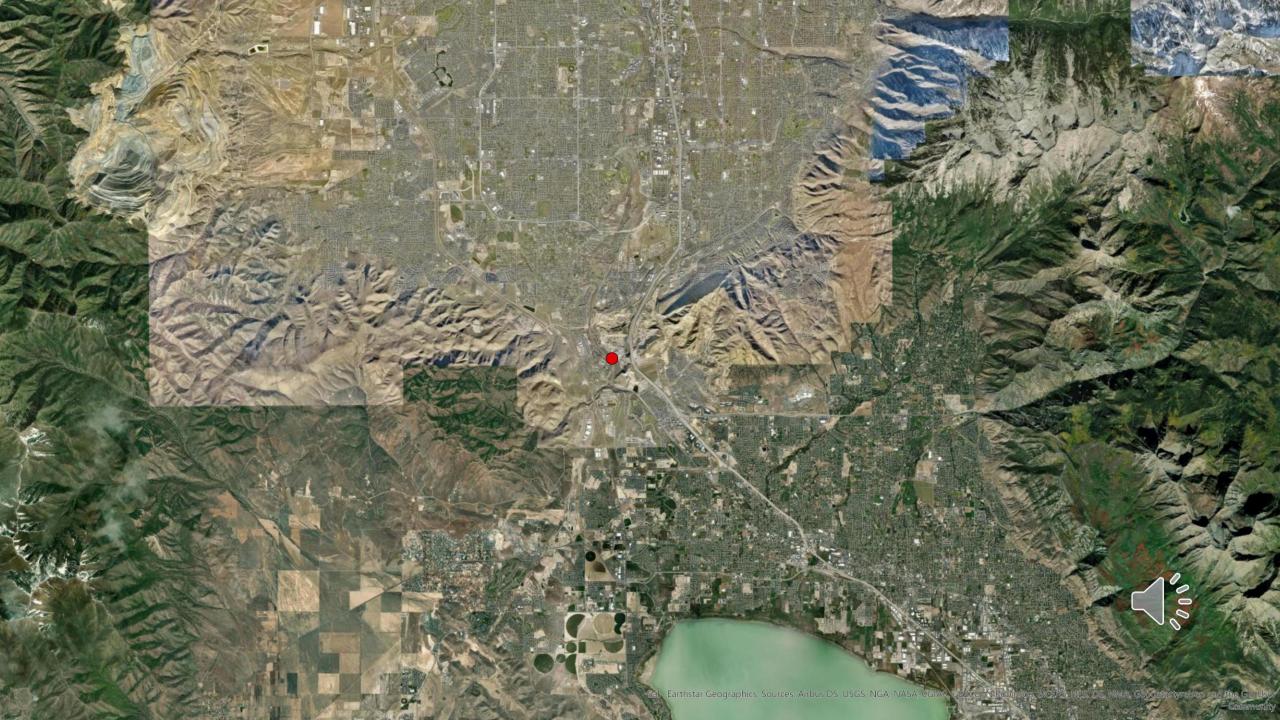
Quantifying Surface Water Influence on Hidden Valley Springs Bluffdale, Utah

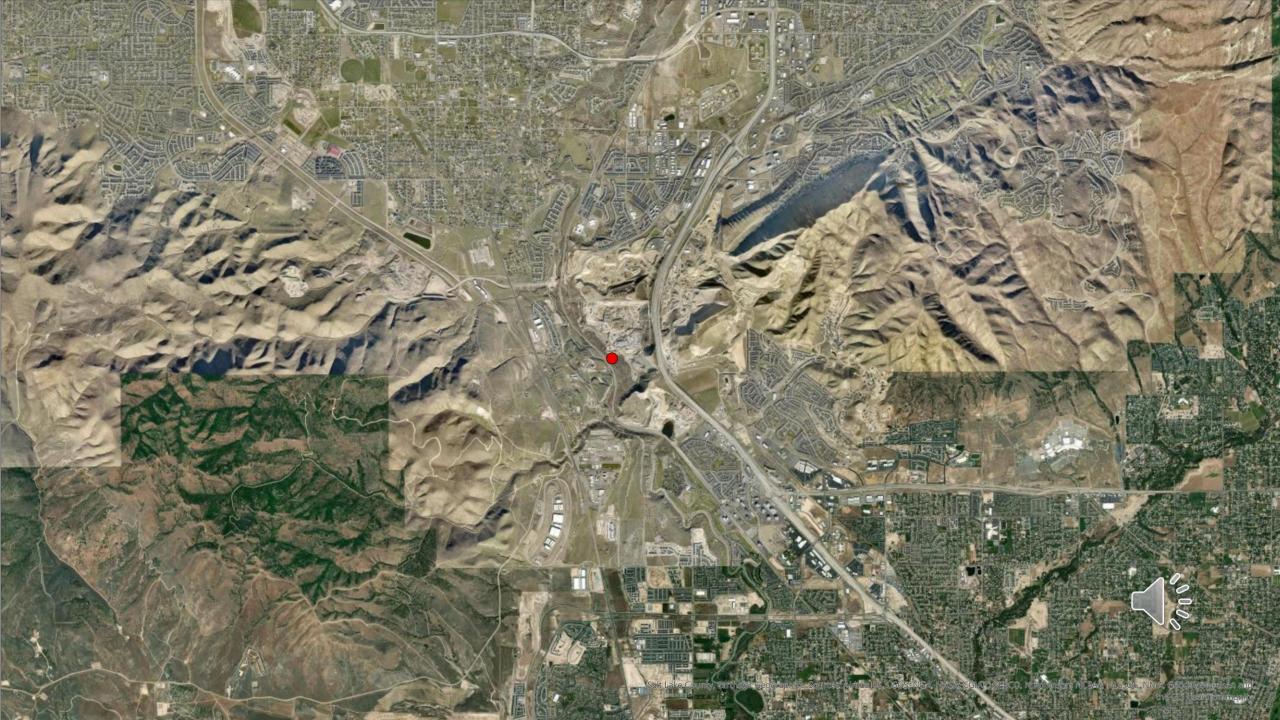
Tyler Yoklavich

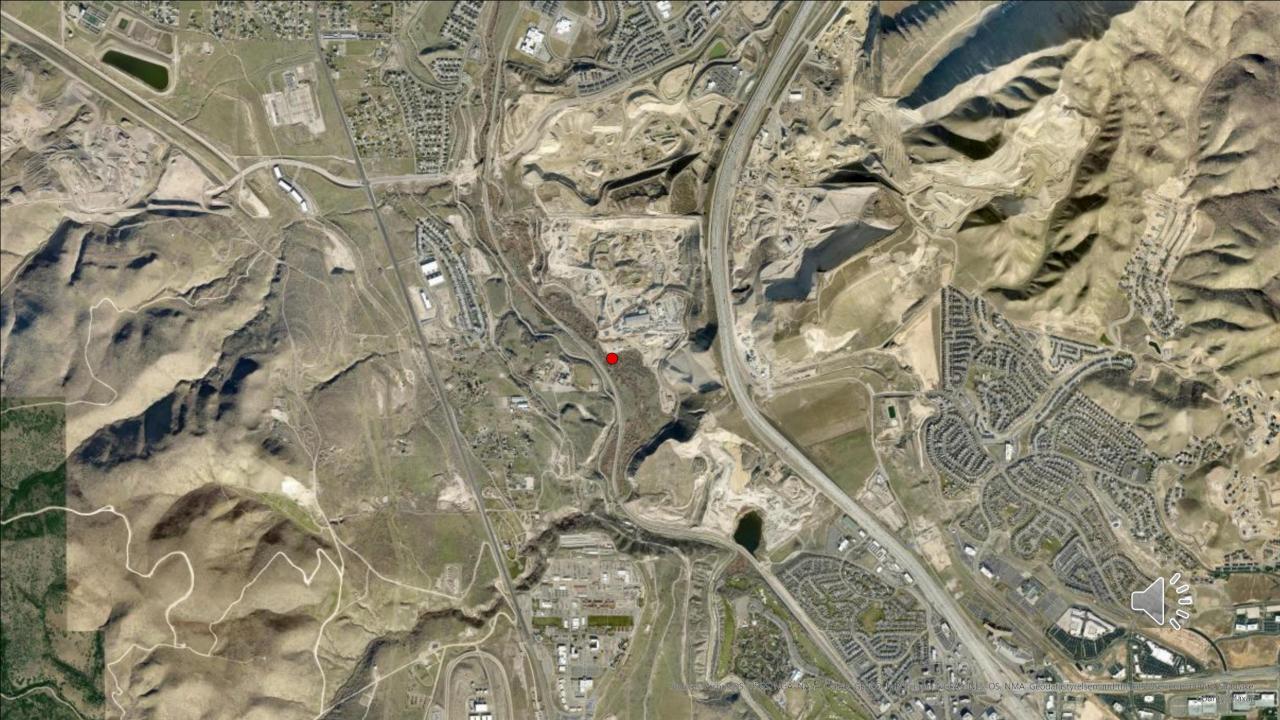
Undergraduate, University of Utah Department of Mines and Earth Sciences



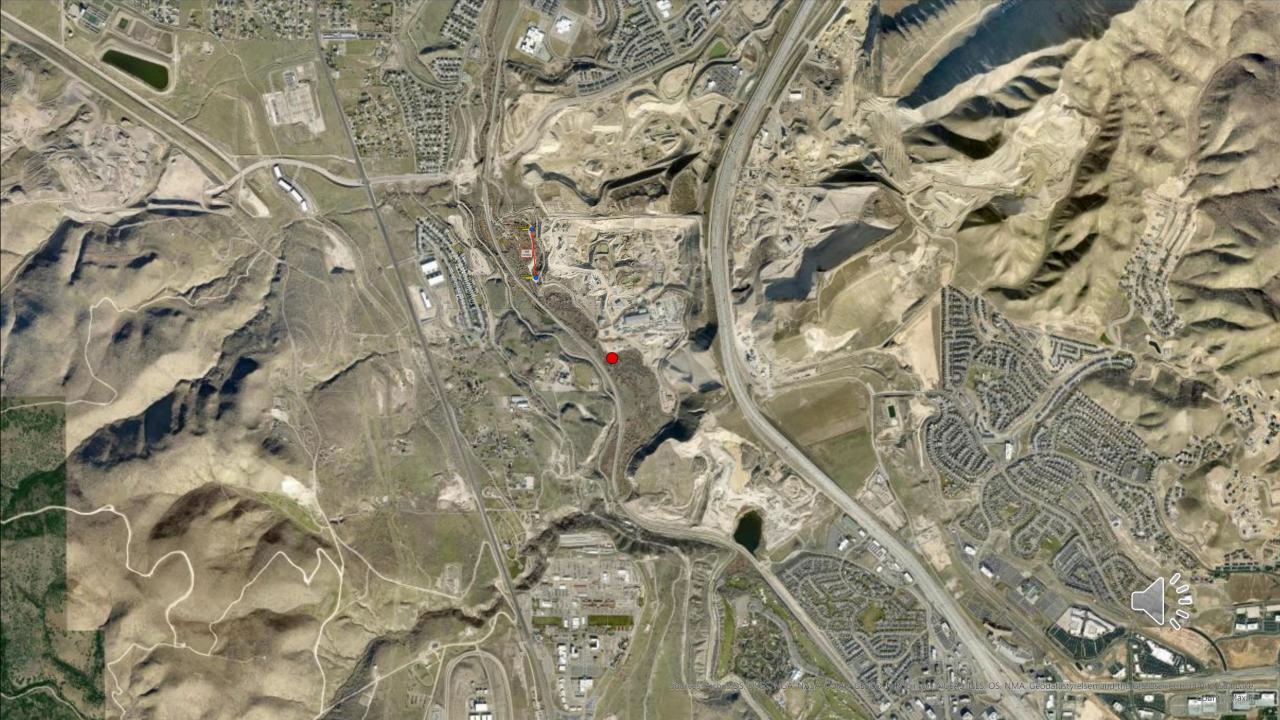


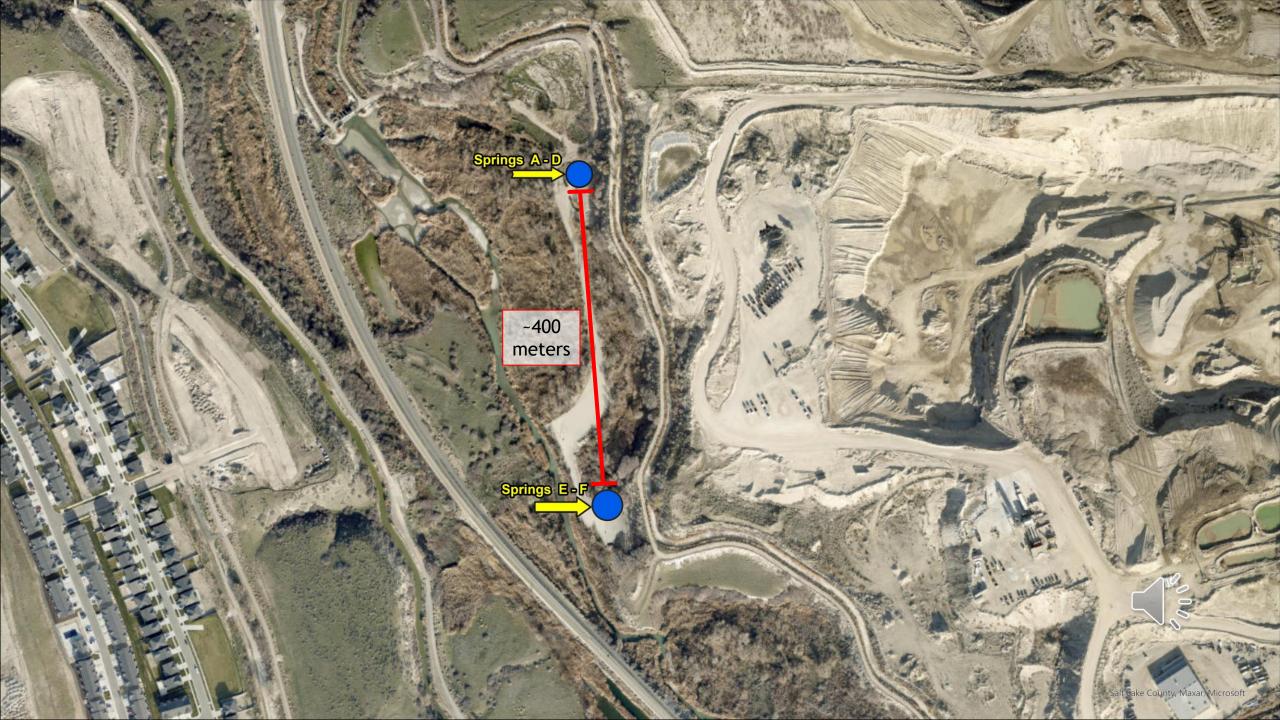












Apparent Age:

- Modern (< 5 years)
 - From ³H, Noble Gases, SF₆, and stable isotopes

Stable Isotope Values:

- Slightly evaporated signature
- Attenuated variation compared to precipitation
 - Suggests potential surface
 water influence
- Plots along same slope as Jordan River

Apparent Age:

~400 meters

prings E-F

- > 50 years
 - From ³H, Noble Gases, SF₆, stable isotopes

Stable Isotope Values:

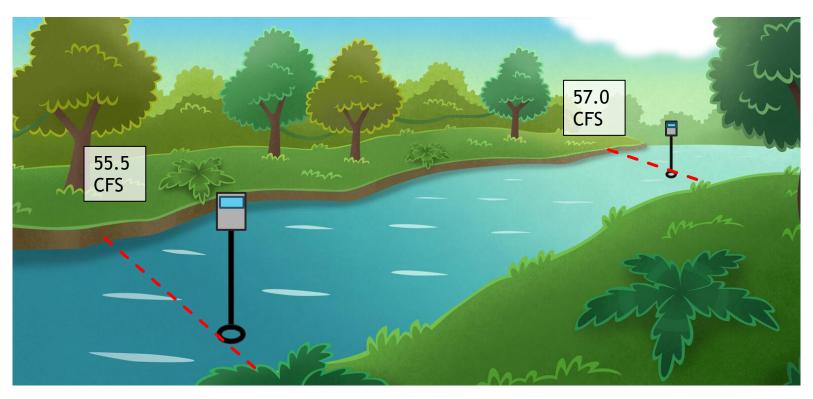
- Plot on Local Meteoric Water Line
 - Suggests precipitation fed recharge with no evaporation

Major Dissolved lons:

Higher TDS, chloride, sodium, sulfate, magnesium, and potassium concentrations than Springs E-F Springs A - D

Geochemical evaluations make it clear that some surface water influence is occurring, but it is difficult to measure how much

Differential Flow Gauging (Seepage Run)



Applied to East Jordan Irrigation Canal:

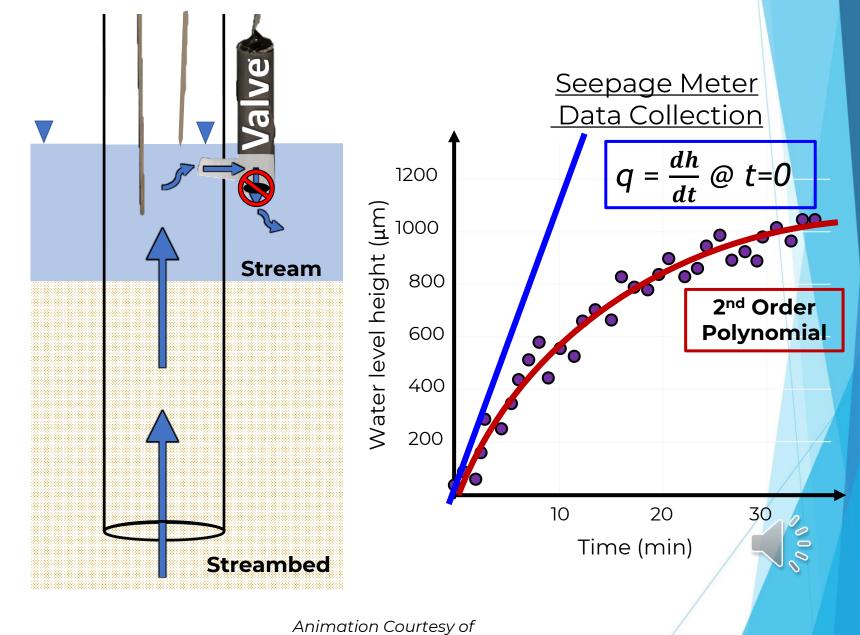
- Reported 1.5 ±2.0 CFS differential
- Springs of interest have mean flow of .9 CFS
- Inconclusive, more precision necessary to constrain estimate

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Automatic Seepage Meter



Solomon et al. (2020)



Dr. Eric Humphrey

Finding a Canal Representative q - Methods



Seepage Measurement Points



- Meters placed at 15 m intervals along centerline of canal
- Measurements taken during 4 nonconsecutive field outings
- Total of 255 individual measurements taken at 21 discrete measurement points
- Each measurement point produced a mean seepage flux (q) in cm/day

Results



- Most measurement points recorded a small amount of mean negative (downward) seepage
- Uncertainty is obtained by calculating the standard error of all flux values recorded at a given point
 - Points with a positive flux may be due to hyporheic flow

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Representative q

- To find a representative seepage flux (q), the average of all measurement point fluxes was calculated
 - Overall uncertainty is obtained by propagating the error associated with each measurement point flux through the averaging of fluxes
- The representative seepage flux (q) was identified to be -.9 ±2.1 cm/day



Meter/	Mean Seepage	Uncertainty +/-
Measurement	Flux, q (cm/d)	(Standard
Point		Error)
R1-1	-0.12	0.96
R1-2	-1.52	0.12
R1-3	-0.78	0.22
R1-4	-1.19	0.11
R1-5	-0.75	0.19
R1-6	3.63	0.51
R1-7	-1.47	0.29
R2-1	-1.15	0.93
R2-2	-0.28	0.51
R2-3	-0.35	0.27
R2-4	1.31	1.14
R2-5	-0.40	0.15
R2-6	-0.34	0.08
R3-2	-0.48	0.06
R3-3	-2.37	0.08
R3-4	-0.46	0.11
R3-5	-3.53	0.07
R3-6	-5.51	0.09
R4-1	0.15	0.19
R4-2	-0.15	0.48
R4-3	-2.62	0.77
Overall:	-0.87	2.08

Estimating a Volumetric Seepage Rate (Q) and Proportion of Spring Flow Due to Canal Seepage

The representative seepage flux (cm/day) was multiplied by the surface area of canal within the springs presumed recharge zone to produce a volumetric seepage rate of $181 \pm 444 \text{ m}^3/\text{day}$ (0.07 ±.18 CFS)

Uncertainty is still high relative to the total seepage, but the precision is much greater than with differential flow gauging (2 CFS error vs .18)

This means that annually, $3.3 \times 10^4 \, \text{m}^3/\text{year seeps}$ from the canal into springs A-D

In 2021, springs A-D had a total flow of 6.4×10^5 m³

Dividing total flow by seepage input indicates that at maximum, ~5% of spring flow may be attributed to seepage losses from the irrigation canal

Limitations/Future Work

Higher than typical uncertainty was recorded during this project

Assumptions made regarding recharge zones and flow paths may not be accurate

All seepage unlikely to discharge at springs

High spatial variability -larger scope may capture more high seepage areas/reduce overall uncertainty

Monitoring wells could establish more well constrained flow paths

Due to these assumptions, we consider 5% to be the maximum possible contribution to spring flow



Summary and Conclusions

- 255 automatic seepage meter measurements were taken along a 400 meter transect of the East Jordan Irrigation Canal
- A representative seepage flux of -.9 ±2.1 cm/day was obtained
- Applying this flux to the area of canal within the springs presumed recharge zone results in a volumetric seepage rate of 181 ± 444 m³/day (0.07 ±.18 CFS), or 3.3x10⁴ m³/year
- A comparison of seepage inputs to total annual spring flow shows that a maximum of 5% of spring flow may be attributed to seepage losses
- Due to these findings, we conclude that the irrigation canal is not a primary source of recharge to Hidden Valley Springs
- Further investigations including more seepage measurements, shallow monitoring well installation, and continued geochemical investigation will help to further constrain recharge sources to Hidden Valley

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